Emissions from miombo woodland and dambo grassland savanna fires

Parikhit Sinha,¹ Peter V. Hobbs,¹ Robert J. Yokelson,² Donald R. Blake,³ Song Gao,^{4,5} and Thomas W. Kirchstetter⁶

Received 8 January 2004; revised 15 March 2004; accepted 31 March 2004; published 10 June 2004.

[1] Airborne measurements of trace gases and particles over and downwind of two prescribed savanna fires in Zambia are described. The measurements include profiles through the smoke plumes of condensation nucleus concentrations and normalized excess mixing ratios of particles and gases, emission factors for 42 trace gases and seven particulate species, and vertical profiles of ambient conditions. The fires were ignited in plots of miombo woodland savanna, the most prevalent savanna type in southern Africa, and dambo grassland savanna, an important enclave of miombo woodland ecosystems. Emission factors for the two fires are combined with measurements of fuel loading, combustion factors, and burned area (derived from satellite burn scar retrievals) to estimate the emissions of trace gases and particles from woodland and grassland savanna fires in Zambia and southern Africa during the dry season (May-October) of 2000. It is estimated that the emissions of CO_2 , CO, total hydrocarbons, nitrogen oxides (NO_x as NO), sulfur dioxide (SO₂), formaldehyde, methyl bromide, total particulate matter, and black carbon from woodland and grassland savanna fires during the dry season of 2000 in southern Africa contributed 12.3%, 12.6%, 5.9%, 10.3%, 7.5%, 24.2%, 2.8%, 17.5%, and 11.1%, respectively, of the average annual emissions from all types of savanna fires worldwide. In 2000 the average annual emissions of methane, ethane, ethene, acetylene, propene, formaldehyde, methanol, and acetic acid from the use of biofuels in Zambia were comparable to or exceeded dry season emissions of these species from woodland and grassland savanna fires in Zambia. INDEX TERMS: 0305 Atmospheric Composition and Structure: Aerosols and particles (0345, 4801); 0312 Atmospheric Composition and Structure: Air/sea constituent fluxes (3339, 4504); 0322 Atmospheric Composition and Structure: Constituent sources and sinks; 0345 Atmospheric Composition and Structure: Pollution-urban and regional (0305); KEYWORDS: trace gases and particles, savanna fires, emissions and emission factors

Citation: Sinha, P., P. V. Hobbs, R. J. Yokelson, D. R. Blake, S. Gao, and T. W. Kirchstetter (2004), Emissions from miombo woodland and dambo grassland savanna fires, *J. Geophys. Res.*, 109, D11305, doi:10.1029/2004JD004521.

1. Introduction

[2] Biomass burning is the primary source of atmospheric pollution in the tropics [*Crutzen and Andreae*, 1990], and it affects large areas of the world through long-range transport [*Fishman et al.*, 1991]. Savanna fires are the largest source of biomass burning emissions worldwide, and tropical Africa contains about two thirds of the world's savanna

Copyright 2004 by the American Geophysical Union. 0148-0227/04/2004JD004521\$09.00

[*Hao and Liu*, 1994]. Savanna burning is a source of a wide variety of compounds that are important in atmospheric chemistry, including carbon monoxide (CO), nitrogen oxides (NO_x), sulfur dioxide (SO₂), nonmethane organic compounds (NMOC), halocarbons, and particles [*Crutzen and Andreae*, 1990; *Sinha et al.*, 2003a]. Smoke aerosols perturb atmospheric radiation through their effects on light extinction and cloud properties [*Reid et al.*, 1998; *Haywood et al.*, 2003]. The photochemical oxidation of CO and NMOC in the presence of NO_x in smoke produces ozone (O₃) [*Radke et al.*, 1978; *Hobbs et al.*, 2003], a key precursor of the hydroxyl radical (OH), which is the primary oxidant in the troposphere.

[3] In August and September 2000 the University of Washington (UW) Cloud and Aerosol Research Group (CARG), with its Convair-580 research aircraft, participated in the Southern African Regional Science Initiative 2000 (SAFARI 2000) field project [*Swap et al.*, 2003]. One of the goals of SAFARI 2000 was to study the emissions and transformations of smoke from biomass burning in various parts of southern Africa. This paper describes measurements

¹Department of Atmospheric Sciences, University of Washington, Seattle, Washington, USA.

²Department of Chemistry, University of Montana, Missoula, Montana, USA.

³Department of Chemistry, University of California, Irvine, California, USA.

⁴Department of Chemistry, University of Washington, Seattle, Washington, USA.

⁵Now at Department of Environmental Science and Engineering, California Institute of Technology, Pasadena, California, USA.

⁶Lawrence Berkeley National Laboratory, Berkeley, California, USA.

of emissions from two prescribed savanna fires that occurred in miombo woodland and dambo grassland plots in Zambia on 1 September and 5 September 2000, respectively. The miombo woodland ecosystem is the most abundant type of savanna in southern Africa and the largest contiguous block of deciduous tropical woodlands and dry forests in the world, covering ~ 2.8 million km² including much of Africa from 4° to 17°S. Dambo grasslands are major enclaves within miombo woodlands, covering up to 40% of the landscape in some areas [International Geosphere-Biosphere Programme, 1997].

[4] Normalized excess mixing ratios are given at various locations over and downwind of the miombo and dambo fires for a number of gaseous species, including carbon dioxide (CO₂). CO, SO₂, dimethyl sulfide (DMS), NO_x, ammonia (NH₃), hydrogen cyanide (HCN), methane (CH₄), NMOC, halocarbons, and particulates. Emission factors are given for these species, and they are combined with measurements of fuel loading, combustion factors, and area burned to estimate their emissions from woodland and grassland savanna fires in Zambia and southern Africa during the dry season of 2000.

2. Sampling Techniques and Instrumentation

[5] Various trace gas and particle species were measured continuously or intermittently aboard the UW Convair-580 research aircraft. A complete list of the instruments aboard the Convair-580 in SAFARI 2000, and the methods of sampling are given in a technical appendix by P. V. Hobbs included by *Sinha et al.* [2003a]. Only the instruments and techniques that provided the measurements presented in this paper are described here.

[6] Continuous measurements of SO₂ were made using a Teco model 43S pulsed-fluorescence analyzer (precision of 7%, detection limit of 0.1 ppbv). The total concentration of particles (CN) in the size range \sim 0.003–3 µm diameter was measured continuously with a TSI 3025A ultrafine condensation particle counter (precision of 10%). The dry aerosol light-scattering coefficient was measured continuously with an MS Electron three-wavelength (0.45, 0.55, and 0.70 µm) nephelometer.

[7] Evacuated electropolished stainless steel canisters were used to sample ambient air using a stainless steel inlet that passed through the aircraft fuselage. For each canister sample, mixing ratios of selected C_2-C_9 nonmethane hydrocarbons (NMHC), methyl halides, DMS, and methyl nitrate (CH₃ONO₂) were determined by gas chromatography (GC) with flame ionization detection (FID), electron capture detection, and mass selective detection. The precision of the measurements was 3%, and the typical NMHC detection limit was 3 parts per trillion by volume. Mixing ratios of CO₂ (precision of 3%), CO (precision of 5%), and CH₄ (precision of 0.1%) in the canisters were determined using three separate GC/FID systems. A detailed description of the analytical procedure for the canister samples is given by *Colman et al.* [2001].

[8] An airborne Fourier transform infrared spectrometer (AFTIR) with a separate and specially coated inlet that directed ram air through a Pyrex multipass cell with an exchange time of 4-5 s was deployed aboard the UW Convair-580. The AFTIR measured CO₂, CO, CH₄, NO,

nitrogen dioxide (NO₂), O₃, ethene, acetylene, formaldehyde, methanol, acetic acid, formic acid, NH₃, and HCN. The AFTIR technique is described by *Yokelson et al.* [2003].

[9] Aerosol samples collected on quartz filters (Pallflex 2500 QAT-UP) were used to determine the concentration of particulate carbon. The quartz filters were baked before use at 800°C for at least 6 hours to remove carbonaceous impurities. They were then analyzed for total carbon content using the evolved gas analysis method described by Novakov [1981, 1982] and Kirchstetter et al. [2003]. Black carbon (BC) concentrations were measured with an optical transmission technique similar to that described by Rosen and Novakov [1983]. This method compares the attenuation of white light through a loaded filter relative to that of a blank filter. The relationship between optical attenuation (ATN) and the BC concentration (in $\mu g \text{ cm}^{-2}$) is given by ATN = σ BC, where ATN = 100 × ln(I₀/I), I₀ and I are the transmitted light intensities through the blank and loaded filters, respectively, and σ is the mass absorption cross section for BC deposited on the quartz filters (in $m^2 g^{-1}$). A value of 20 m² g⁻¹ was used for the mass absorption cross section [Gundel et al., 1984], which is consistent with the calibration factor of a commercial aethalometer that employs the same optical transmission method for measuring BC concentrations [Bodhaine, 1995]. Further discussion of this optical transmission method, including uncertainties and comparison with other methods, is given by *Kirchstetter* et al. [2003] and Sinha et al. [2003b].

[10] Particles were also collected on Teflon filters (Gelman Sciences Teflon membrane, 2.0 µm pore size). The Teflon filters were weighed before and after particle sampling in a relative humidity (RH)- and temperature (T)controlled chamber (RH = 40%, T = 293 K) to determine the masses of dry particulate matter (PM) collected on the filters. From control and field blank filters the uncertainty of the total dry PM measured with these filters was estimated to be $\pm 6 \mu g$. After gravimetric analysis the Teflon filters were extracted in deionized water (high-performance liquid chromatography grade) and analyzed using a standard ion chromatography system (Dionex DX 500). This analysis yielded mass concentrations of three inorganic ions (nitrate, sulfate, and chloride) with a precision of 5%. An inductively coupled plasma-atomic emission spectrometer (Jarrell Ash 955) was used to measure the mass concentration of the potassium ion to a precision of 4%. Further details on the Teflon filter analyses of aerosol compounds are given by Gao et al. [2003].

[11] When sampling smoke plumes we employed a "grab bag" technique to obtain samples for the filters and SO₂. The grab bag consisted of a 2.5 m³ electrically conducting plastic (Velostat) bag that could be filled with a sample of smoke in 12 s when exposed to ram air. The grab bag system had an aerosol cutoff diameter of about 4 μ m; larger particles were lost in the inlet and on the walls of the grab bag.

3. Definitions and Calculations of Normalized Excess Mixing Ratios, Emission Factors, and Combustion Efficiencies

[12] A normalized excess mixing ratio of species X is obtained by dividing the excess (i.e., above background) mixing ratio of X by the excess mixing ratio of a simulta-

neously measured reference gas (such as CO_2 or CO). For example, a normalized excess mixing ratio (NEMR) of species X relative to CO is

$$\operatorname{NEMR}(X)_{\rm CO} = \frac{[\Delta X]}{[\Delta {\rm CO}]}.$$
 (1)

[13] Emission factors for gases and particles were calculated using the carbon balance mass method [*Radke et al.*, 1988]. The underlying premise of this method is that all of the carbon combusted in a fire and released to the atmosphere is emitted into the smoke plume as CO₂, CO, CH₄, NMOC, and particulate carbon (PC). The emission factor (EF) of a species X is the ratio of the excess mass concentration [ΔX] of X to the excess total carbon concentration [ΔC] in the plume:

$$\mathrm{EF}(X) = \frac{[\Delta X]}{[\Delta C]_{\mathrm{CO}_2} + [\Delta C]_{\mathrm{CO}} + [\Delta C]_{\mathrm{CH}_4} + [\Delta C]_{\mathrm{NMOC}} + [\Delta C]_{\mathrm{PC}}}.$$
 (2)

The emission factor is expressed in units of grams of X emitted per gram of carbon burned. To convert this emission factor to units of grams of X emitted per gram of fuel burned, EF is multiplied by the mass fraction of carbon in the fuel. Typically, the carbon content of biomass fuels varies from 45 to 55% [*Susott et al.*, 1996]; in this paper we assume that it is 50%.

[14] The combustion efficiency (CE) is the ratio of carbon emitted as CO_2 to the total carbon emitted [*Ward and Hardy*, 1991],

$$CE = \frac{[\Delta C]_{CO_2}}{[\Delta C]_{CO_2} + [\Delta C]_{CO} + [\Delta C]_{OC} + [\Delta C]_{PC}},$$
(3)

where the subscript OC indicates total organic compounds (i.e., methane and NMOC). Thus CE is the fraction of fuel carbon emitted that is completely oxidized to CO₂. Combustion efficiency is a useful way to quantify the relative amounts of flaming and smoldering combustion: When CE < 1, the fire emissions were produced by a mixture of flaming and smoldering combustion; when CE > 90%, more than 50% of the emissions were produced by flaming combustion; and when CE < 90%, more than 50% of the emissions were produced by *glaming combustion*; and when CE < 90%, more than 50% of the emissions were produced by smoldering combustion [*Ward and Hardy*, 1991].

[15] Although CE is a useful quantity for fire models, it is often difficult to measure all of the individual carbon species in the emissions from a fire. Therefore, in this study we use the modified combustion efficiency (MCE) to describe the relative amounts of flaming and smoldering combustion [*Ward and Hao*, 1992], which is defined as

$$MCE = \frac{[\Delta C]_{CO_2}}{[\Delta C]_{CO_2} + [\Delta C]_{CO}}.$$
 (4)

Since OC and PC are emitted in relatively small quantities relative to CO_2 and CO, the difference between CE and MCE is typically only a few percent.

4. Vegetation and Fires

[16] The prescribed fire on 1 September 2000 was ignited at 0846 UTC (local time is UTC plus 2 hours) at 14.82° S, 24.48°E near Kaoma, Zambia. The fire burned for ~90 min



Figure 1. Photograph of a dambo grassland fire in Zambia taken at 1157 UTC (8 min after ignition) on 5 September 2000 from the University of Washington Convair-580 research aircraft. (Photograph courtesy of P. V. Hobbs.)

over 95 ha (J. M. C. Pereira et al., unpublished manuscript, 2002) The vegetation type, as classified by *White* [1981, 1983], was "wetter Zambezian miombo woodland." This ecosystem, which we will refer to as "miombo woodland," is dominated by the tree genera *Brachystegia*, *Julbernardia*, and *Isoberlinia* [*White*, 1981, 1983]. The vegetation combusted in the miombo woodland fire considered here consisted of standing grass, leaf and twig litter, large fallen dead wood, and live leaves. Fuel loading, combustion completeness, fire intensity, and rate of spread of this fire are given by J. M. C. Pereira et al. (unpublished manuscript, 2002).

[17] The prescribed fire on 5 September 2000 was ignited at 1149 UTC at 14.81°S, 24.46°E, also near Kaoma, Zambia. The fire burned for \sim 50 min over 70 ha (J. M. C. Pereira et al., unpublished manuscript, 2002). The vegetation type, as classified by White [1981, 1983], was "edaphic grassland mosaics with semiaquatic vegetation." This ecosystem, which we will refer to as "dambo grassland," occupies seasonally waterlogged shallow valley depressions with a medium-dense, uniform grass mat dominated by the grass species Loudetia simplex [White, 1981, 1983]. The vegetation combusted in the dambo grassland fire considered here consisted exclusively of standing grass. Other biomass parameters for this fire are given by J. M. C. Pereira et al. (unpublished manuscript, 2002). Figure 1 shows an image of the dambo fire taken at 1157 UTC, 8 min after ignition.

5. Results

5.1. Background Conditions

[18] The miombo and dambo fires took place in different background meteorological conditions. In the first week of September 2000 the passage of a cold front originating in the subtropics induced northwesterly flow throughout southern Africa. Heavy smoke from intense biomass burning regions in Angola was transported over Zambia and other southern African nations toward the Indian Ocean in an event that has been termed the "river of smoke" [*Annegarn et al.*, 2002]. On 1 September the river of smoke was just beginning, but by 5 September it was well



Figure 2. Vertical profiles of temperature, relative humidity, SO₂, O₃, CN, and the light-scattering coefficient (at $\lambda = 0.55 \ \mu$ m) measured upwind of the miombo woodland fire (circles) and dambo grassland fire (crosses) from 0650 to 0856 UTC on 1 September 2000 and from 1054 to 1150 UTC on 5 September 2000, respectively.

underway. Consequently, background conditions for the miombo fire on 1 September were generally less polluted than for the dambo grassland fire on 5 September.

[19] Background vertical profiles of temperature, relative humidity, SO₂, O₃, CN, and light-scattering coefficient (at $\lambda = 0.55 \ \mu$ m) for the miombo and dambo fires are shown in

Figure 2. In both cases, temperature lapse rates were generally $8^{\circ}-9^{\circ}$ C km⁻¹, resulting in conditionally unstable atmospheres. For the miombo fire, stable layers were present at ~1.75 and ~3.25 km, which inhibited vertical transport of the smoke. For the dambo fire the atmosphere was well mixed up to at least 4 km.

[20] Ozone concentrations upwind of the dambo fire were \sim 20–30 ppbv higher than those upwind of the miombo fire, reflecting the transport of aged biomass burning emissions within the river of smoke. This is supported by background CO concentrations measured by AFTIR for the miombo (\sim 300 ppbv) and dambo (\sim 500 ppbv) fires, which indicate a more polluted background for the dambo fire than the miombo. Also, upwind of the fires, the light-scattering coefficient ranged from 4×10^{-5} to 2×10^{-4} , with higher values upwind of the dambo fire than upwind of the miombo fire. On the other hand, background CN and SO₂ concentrations were similar for both cases, ranging from $\sim 5 \times 10^3$ to 1×10^4 cm⁻³ and from 1 to 2.5 ppbv, respectively, near the surface. Despite similar CN profiles the difference in the light-scattering coefficient between the two cases suggests different aerosol size distributions with more combustion-generated submicron particles upwind of the dambo grassland fire than the miombo woodland fire. The relative humidity was low (<35%) upwind of both fires. Therefore the effect of RH on the aerosol light-scattering coefficient was small [Magi and Hobbs, 2003]. The average surface wind speed upwind of the miombo and dambo fires was 5.6 ± 2.1 m s⁻¹ from the north and 6.3 ± 2.2 m s⁻¹ from the northeast, respectively.

5.2. Plume Dimensions

[21] Horizontal cross sections of the plumes of smoke from the miombo and dambo fires are shown in Figure 3. The width and length of the miombo plume were $\sim 3-6$ and \sim 14 km, respectively, and the width and length of the dambo plume were $\sim 2-3$ and ~ 5 km, respectively. The smaller plume dimensions of the dambo fire reflect in part the smaller area burned (70 ha) compared to the area burned by the miombo fire (95 ha). Also, the duration of the dambo fire $(\sim 50 \text{ min})$ was less than that of the miombo fire $(\sim 90 \text{ min})$. Surface wind speeds were similar in both cases ($\sim 6 \text{ m s}^{-1}$). Since the combusted vegetation in the dambo fire consisted exclusively of standing grass, it burned more rapidly than the miombo fire, which consumed both standing grass and less flammable fuel such as large, fallen dead wood. In both fires, CN concentrations peaked at $\sim 10^6$ cm⁻³ over the fire and decreased to nearly background concentrations of $\sim 10^4$ cm⁻³ several kilometers downwind of the fires. The high background concentrations of CN limited the distances downwind that the smoke plumes could be detected.

5.3. Plume Samples

[22] Smoke samples from both fires were obtained in a series of passes perpendicular to the plume axes at various distances and altitudes downwind of the fires. Table 1 summarizes the time, location, age, and various state parameters for these samples. Ground level was \sim 1100 m, and most smoke samples were obtained <300 m above ground level. The age of each smoke sample was estimated by dividing the downwind distance of the sample by the average wind speed. The MCE of each sample indicates the stage of combustion during which the sample was released, with values >0.90 indicating predominantly flaming combustion.

5.4. Normalized Excess Mixing Ratios

[23] Table 2 shows the excess mixing ratios of 49 trace gas and particle species normalized with respect to excess



Figure 3. Condensation nucleus concentrations across the widths of (a) miombo woodland and (b) dambo grassland smoke plumes at three distances downwind of each of the fires.

CO for samples 1-8 from the miombo woodland fire and samples a-f from the dambo grassland fire described in Table 1. The data set is unique in providing plume measurements of recently emitted smoke of known ages (~1-40 min and $\sim 1-4$ min for the miombo woodland and dambo grassland savanna fire, respectively (Table 1)) for a large number of trace gas and particle species, including carbon dioxide, carbon monoxide, sulfur dioxide, nitrogen oxides, methane, ammonia, dimethyl sulfide, nonmethane organic compounds, halocarbons, gaseous organic acids, aerosol ionic components, carbonaceous aerosols, and condensation nuclei. Additionally, because the fires in this study were prescribed, the combustion processes at the ground were well characterized (section 4). A discussion of the variations in the measured emissions from southern African savanna fires is given by Sinha et al. [2003a] and Yokelson et al. [2003], and the subsequent evolution of the emissions is discussed by Hobbs et al. [2003].

[24] The downwind transformations of chemical species in the plume of a 1000 ha savanna fire in South Africa

Sample ^a	Time, UTC	Latitude, °S	Longitude, °E	Distance From Fire, m	Altitude, m	Age of Smoke, min	Temperature, °C	Pressure, hPa	Relative Humidity, %	Modified Combustion Efficiency
1	0857	14.798	24.487	2,549	1,340	7.5	23	870	17.5	0.935
2	0902	14.816	24.478	517	1,325	1.5	24	871	17.0	0.917
3	0909	14.826	24.477	744	1,257	2.2	24.8	878	17.0	0.944
4	0919	14.829	24.482	1,019	1,330	3.0	23.8	871	16.7	0.942
5	0927	14.817	24.461	2,073	1,905	6.1	18.3	814	25.3	0.952
6	0939	14.822	24.364	12,495	1,976	36.8	18.1	810	25.1	0.913
7	0947	14.806	24.349	14,150	1,959	41.7	18.8	811	18.9	
8	1006	14.814	24.444	4,009	2,055	11.8	18.1	803	21.0	0.985
а	1155	14.812	24.457	374	1,292	1.0	30.1	871	10.9	0.923
b	1157	14.814	24.455	750	1,339	2.0	29.7	867	10.3	0.960
с	1202	14.815	24.454	823	1,340	2.2	30.4	865	8.8	0.977
d	1205	14.814	24.453	912	1,323	2.4	32.2	868	9.7	0.976
e	1209	14.818	24.451	1,259	1,343	3.3	29.4	865	8.7	0.966
f	1233	14.814	24.448	1,346	1,347	3.6	28.8	863	7.4	0.961

Table 1. Samples of Smoke From the Miombo Woodland Fire on 1 September 2000 and the Dambo Grassland Fire on 5 September 2000

^aSamples 1-8 are from the miombo woodland fire; samples a-f are from the dambo grassland fire.

(the Timbavati fire) have been described by *Hobbs et al.* [2003]. The Timbavati fire occurred in the semiarid south Zambezian undifferentiated woodland savanna of the South African lowveld. Various alkenes and NO₂ were observed to be depleted in smoke aged about 40-45 min downwind of the Timbavati fire, while other species such as gaseous acetic acid, ozone, and particulate nitrate were enhanced in the aged smoke.

[25] Since the areal extents of the two Zambian fires were an order of magnitude less than those of the Timbavati fire, fewer downwind measurements were obtained in these fires. Nevertheless, downwind enhancements of gaseous acetic acid and particulate nitrate were observed in the Zambian fires (Table 2). Combining the normalized excess mixing ratios from the miombo woodland and dambo grassland fires with those from the Timbavati fire provides linear estimates of gaseous acetic acid and particulate nitrate production in aged smoke downwind of southern African savanna fires (Figure 4).

[26] In addition to nitrate particles the normalized excess mixing ratios of total particulate matter, organic carbon, chloride, and sulfate were enhanced by 60-300% in smoke aged ~ 37 min relative to smoke aged ~ 1.5 min for the miombo woodland fire (Table 2). Oxidation of NMOC, NO_x, and SO₂ and subsequent gas-particle conversion has been shown to enhance particle distributions in aged smoke downwind of savanna fires [*Gao et al.*, 2003; *Hobbs et al.*, 2003].

[27] Normalized excess mixing ratios of four species (CO₂, CH₄, ethene, and acetylene) were obtained by two techniques (GC/C and AFTIR) in the smoke from the two Zambian fires. Normalized excess mixing ratios from these techniques for samples obtained at approximately the same time are compared in Figure 5. The slope of the correlation is 1.03 ± 0.12 with a correlation coefficient (r^2) of 0.88, indicating reasonable agreement between these two independent measurement techniques, except for ethene for which emission factors obtained by AFTIR were about a factor of 2 greater than those obtained by GC/C. It should be noted, however, that these measurements were not made on exactly the same smoke samples since the two measurement techniques

had different sampling inlets and slightly different sampling times.

5.5. Emission Factors and Estimates of Regional Emissions

[28] Estimates of fuel consumed by woodland and grassland savanna fires in southern Africa are presented in Table 3. Combining these fuel consumption estimates with emission factors for 42 trace gas and seven particulate species from the miombo woodland and dambo grassland fires (Table 4), we can obtain estimates of the emissions of these trace gases and particles from woodland and grassland savanna fires in southern Africa during the dry season (May–October) of 2000 (Table 4).

[29] The fuel consumption values in Table 3 were calculated by multiplying burned area by the amount of fuel biomass above ground and by a combustion factor (CF), defined as the fraction of the available fuel biomass above ground combusted in a fire. Burned areas by month and country in southern Africa in 2000 for two broad vegetation types, woodland savanna and grassland savanna, were obtained from SPOT satellite burned area retrievals (Global Burnt Area 2000 Project, Joint Research Centre, Ispra, Italy, available at www.grid.unep.ch/activities/earlywarning/ preview/ims/gba). The burned-area estimates used in our study differ from those reported by Silva et al. [2003] because of updates in the SPOT burned-area algorithm. The amount of biomass fuel loading above ground in plots of miombo woodland and dambo grassland in Zambia was measured by Hoffa et al. [1999], Shea et al. [1996] and J. M. C. Pereira et al. (unpublished manuscript, 2002) with average values of 0.87 ± 0.30 kg m⁻² and $0.34 \pm$ 0.10 kg m^{-2} for miombo woodland and dambo grassland, respectively. Since the moisture content of vegetation decreases as the dry season progresses, the CF increases from May to October in southern Africa [Hoffa et al., 1999]. The fuel consumption measurements indicate that for miombo woodland fires the CF increases fairly linearly from a value of ~ 0.01 at the beginning of the dry season to ~ 0.89 near the end of the dry season. For dambo grassland fires the CF remains at 0.7 ± 0.2 throughout the dry season.

$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	6							
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		8 /.	а	þ	с	q	е	f
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		64,800		24,900			28,800	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	10,500	1 74 1	12,100	22,800	43,200	40,200	275	24,600
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	55.6	T • L /	26.3	27.8	44.2	40.9	C:11	17.5
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$			6.4	7.8	30.1	41.4		22.9
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	r 44		28.4	40.9	48.1	40.0		51.3
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	C: 41		50	8	13.6	12.6		
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		0.0019	2	0.0027		i	0.0012	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		4.4 11.1		1.2			6.1	6.9
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	00		6 0	0.01		12.0		0 1 0
	0.6	0.0046	0.0-	-13.4	C.77—	0.01-	0 0011	0.10-
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		0.41		0.23			0.31	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$		0.0012		0.0001			0.0004	
		3.98		1.66			1.83	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		3.41		8.91			9.35	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			16.82	19.48	26.06	24.89		14.84
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		0.73		0.24			0.25	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		0.56		1.53			1.72	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		2.76		4.19			3.85	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			4.26	6.12	8.69	7.82		
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		0.02		0.01			0.01	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		0.11		c0.0			0.05	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$		0.08		0.04			0.05	
0.006 0.005 0.03 0.005 1.06 0.63 0.79 0.37 0.79 0.37 0.37 0.37 0.37 0.37 0.35 0.37 0.36 0.04 0.05 0.04 0.36 0.04 0.36 0.04		0000		0.03			0.04	
0.03 1.06 0.79 0.79 0.79 0.37 0.36				0.00			0.00	
0.00 0.79 0.79 0.37 0.79 0.37 0.36 0.37 0.37 0.37 0.37 0.37 0.37 0.37 0.36 0.37 0.37 0.37 0.36 0.36 0.36 0.37 0.36 0.56		0.03		0.00			0.00	
0.79 0.37 0.79 0.37 0.79 0.37 0.80 0.37 0.35 0.36 0.78 0.36 0.78 0.36 0.78 0.36 0.78 0.36		C0.0		0.02			0.34	
0.79 0.37 0.79 0.37 0.80 0.37 0.78 0.36 0.78 0.36 0.04 0.25 0.78 0.36				0.02			0.02	
0.79 0.37 0.79 0.37 0.80 0.37 0.78 0.36 0.04 0.78 0.25 0.25 0.78 0.36								
0.79 0.37 0.80 0.37 0.78 0.36 0.25 0.78 0.25 0.78 0.36				0.01			0.02	
0.80 0.37 0.78 0.36 0.05 0.04 0.25 0.25 0.78 0.36 0.78 0.36				0.01			0.01	
0.78 0.36 0.05 0.04 0.35 0.25 0.78 0.36 0.78 0.36				0.01			0.02	
0.78 0.36 0.05 0.04 0.25 0.25 0.78 0.36 0.78 0.36								
0.05 0.04 0.35 0.25 0.78 0.36 0.07 0.02				0.01			0.01	
0.35 0.25 0.78 0.36 0.02 0.02		0.02		0.06			0.05	
		1.13		0.32			0.39	
0.02 0.07								
20:0				0.03			0.02	
0.64 0.65		0.66		0.76			0.77	
0.49 0.47 0.47 0.47 0.40 0.47 0.47 0.61 0.47 0.01		0.34	67	0.55	63	146	0.41	
13.7 15.0 12.1 11.6 13.9			10.4	1.9.1	0.2 9.1	14.0 8.7		17.5
22.8 17.0 13.9 8.4 23.9	63.6		18.2	18.1	10.9	10.0		

D11305

SINHA ET AL.: EMISSIONS FROM SAVANNA FIRES

D11305

	Measurement							San	ple						
Species	Technique ^b	1	2	3	4	5	6	7	8	а	þ	с	q	e	f
Total	F/GB		172.6					438.8			96.1				
particulates															
Organic carbon	F/GB		41.0					59.2			14.5				
Black carbon	F/GB		6.3					7.0			5.9				
Chloride	F/GB		3.5					5.1			0.4				
Nitrate	F/GB		2.6					7.6			2.0				
Sulfate	F/GB		1.8					5.6			2.8				
Potassium	F/GB		5.1					3.0			3.0				
CN°	TSI 3025A	$8.2 imes10^5$	1.2×10^{6}	4.4×10^{5}	8.3×10^{4}	1.3×10^5 ($6.5 imes 10^4$ 2	2.3×10^{4}	1.9×10^{5}	$1.9 imes 10^5$	$2.0 imes 10^5$	1.7×10^5	$3.0 imes10^5$	$3.2 imes 10^5$	3.3×10
^a Trace gas rat ^b Abbreviation:	ios are given in s are GC/C, gas	ppbv ppmv chromatogi	⁷⁻¹ of CO; part	ticle species ra sters; AFTIR,	atios are given airborne Four	i in μg m ⁻³] ier transform	ppmv ⁻¹ of C infrared spe	O. sctroscopy; a	and F/GB, fil	ters via grab	bag.				
^c Units are nur	nber ppmv ⁻¹ o.	f CO.													



Figure 4. Excess mixing ratios of (a) gaseous acetic acid and (b) particulate nitrate normalized with respect to CO as a function of age of the smoke from the miombo woodland, dambo grassland, and the Timbavati fires.

[30] In Figure 6 the fuel consumption values in Table 3 are summed and plotted by month for the dry season of 2000. These results suggest that the biomass combusted in woodland and grassland savanna fires in southern Africa in 2000 peaked in July and September at a value of ~90 Tg. The July peak was primarily due to burning in Angola and the Democratic Republic of Congo, whereas the September peak was largely attributable to burning in Angola, Mozambique, and Zambia (Table 3). Eck et al. [2003] report monthly means of aerosol optical depth (at a wavelength of 500 nm) from 1995 to 2000 at Mongu, Zambia. This study showed that as the dry season progresses, aerosol optical depth increases and peaks at a value of ~ 0.65 in September, the month of highest biomass combustion in Zambia in 2000 (18,372 Gg, Table 3). We also note that Swap et al. [2003] proposed that biomass combustion in southern Africa during the 2000 dry season was above average because of aboveaverage rainfall in the preceding wet season. Above-average rainfall would enhance fuel loading, but changes in com-

.



Figure 5. Comparisons of excess mixing ratios of CO_2 , CH_4 , ethene, and acetylene, normalized with respect to CO, for the dambo grassland and miombo woodland fire plumes determined from gas chromatography measurements on canister samples (GC/C) and from airborne Fourier transform infrared spectrometer measurements (AFTIR).

bustion factors and burned area would also have to be considered to determine interannual variability in biomass combustion. A study of interannual variability of biomass burning emissions constrained by satellite observations [*Duncan et al.*, 2003] found that emissions of CO from biomass burning in southern Africa in 2000 were not above average with respect to the 1979–2000 study period.

[31] In Table 4 we give estimates of emissions from all woodland and grassland savanna fires in southern Africa from May-October of 2000, the bulk of the southern African dry season. These estimates can be compared with the average annual global emissions from savanna burning worldwide estimated by Andreae and Merlet [2001] (Table 4) to obtain rough comparisons of the contributions of emissions from woodland and grassland savanna fires in the southern African dry season to the annual emissions from savanna burning worldwide. Andreae and Merlet estimated emissions assuming that a total of 3160 Tg of biomass is burned annually in savanna fires worldwide. However, the uncertainty in this total was large enough to make a defensible error analysis impossible. In Table 3 we estimate that \sim 367 Tg of biomass were burned in woodland and grassland savanna fires in southern Africa during the dry season (May-October) of 2000. This is within the (large) range of previous estimates for this region (90-2719 Tg [Scholes et al., 1996]) and would account for $\sim 12\%$ of the biomass that Andreae and Merlet [2001] estimated is burned annually by savanna fires worldwide.

[32] Comparing species emissions from woodland and grassland savanna fires in southern Africa in the dry season of 2000 with those from global annual savanna burning [*Andreae and Merlet*, 2001], we find that the woodland and grassland savanna fires in southern Africa accounted for $\sim 12.3\%$, 12.6%, 5.9%, 10.3%, 4.0%, and 7.5% of the annual emissions of CO₂, CO, hydrocarbons (CH₄ + NMHC), NO_x (as NO), NH₃, and SO₂, respectively, from all types of savanna fires worldwide. For total particulate matter, black carbon, organic carbon, and potassium the corresponding percentages are $\sim 17.5\%$, 11.1%, 9.9%, and 12.2%, respectively. For oxygenated species, such as formaldehyde, acetic acid, formic acid, and methanol, the per-

Table 3. Biomass Combusted by Woodland and Grassland Savanna Fires in Southern Africa During the Dry Season of 2000^a

Country	Ecosystem	May	June	July	August	September	October	May-October
Angola	grassland	89	1,260	2,015	1,182	594	93	5,232
Angola	woodland	309	12,904	34,092	28165	18,182	3,460	97,112
Botswana	grassland	24	174	10	29	104	205	546
Botswana	woodland	89	1,878	171	733	3,334	8,052	14,257
Congo ^b	grassland	671	2,314	2,244	621	93	12	5,956
Congo ^b	woodland	1,581	16,188	25,935	10,098	1,953	311	56,066
Lesotho	grassland	14	15	10	32	13	1	86
Lesotho	woodland	11	36	40	177	91	6	362
Malawi	grassland	2	2	34	15	43	16	111
Malawi	woodland	5	13	458	279	1,047	460	2,263
Mozambique	grassland	2	6	145	301	294	181	928
Mozambique	woodland	12	125	4,975	14,530	18,252	13,743	51,637
Namibia	grassland	45	17	18	38	120	34	272
Namibia	woodland	309	344	613	1,788	7,326	2,547	12,927
South Africa	grassland	189	1,275	709	753	1,471	332	4,729
South Africa	woodland	118	2,356	2,164	3,236	8,126	2,241	18,242
Swaziland	grassland	<1	<1	8	2	6	<1	17
Swaziland	woodland	<1	2	59	26	88	3	178
Tanzania	grassland	166	983	834	493	348	366	3,189
Tanzania	woodland	342	5,989	8,390	6,982	6,328	8,139	36,170
Zambia	grassland	39	259	561	294	609	222	1,985
Zambia	woodland	134	2,612	9,344	6,899	18,372	8,181	45,542
Zimbabwe	grassland	2	3	29	71	220	136	461
Zimbabwe	woodland	4	19	306	1,050	4,159	3,131	8,668

^aSee section 5.5 for details of how the quantity of biomass combusted was derived in this study. "Southern Africa" is defined as the African continent below 4°S, excluding Madagascar. Values are given in Gg.

^bDemocratic Republic of Congo

Species	Emission Factor for the Miombo Fire ^b	Emission Factor for the Dambo Fire ^b	Emissions From Woodland Savanna Fires in Southern Africa May–October 2000	Emissions From Grassland Savanna Fires in Southern Africa May–October 2000	Annual Emissions From Savanna Fires Worldwide ^c	Emissions From Woodland and Grassland Savanna Fires in Zambia May–October 2000	Annual Emissions From Use of Biofuels in Zambia ^d
CO ₂	1,705	1,759	585,500	41,400	5,096,000	81,100	13,100
CO	73	42	25,100	988	206,000	3,410	970
CH ₄	1.4	0.5	481	12	7,400	65	109
NMHC	1.7	1.1	584	26	10,700	80	
NO_x (as NO)	3.5	2.4	1,202	56	12,200	164	14
NH ₃	0.4		137		3,400	18	10
HCN	0.37	0.31	127	7.3	90	17	
MeONO ₂	0.00028	0.00023	0.096	0.005		0.01	
SO ₂	0.22	0.3	75.6	7.1	1,100	10.6	
DMS	0.00049	0.00016	0.17	0.004		0.02	
CH ₃ Br	0.00048	0.0001	0.16	0.002	6	0.02	
CH ₃ Cl	0.029	0.02	9.96	0.47	240	1.36	
CH ₃ I	0.000084	0.000055	0.03	0.001	6	0.004	
Ethane	0.19	0.078	65.3	1.8		8.8	20
Ethene	0.43	0.38	148	8.9		20.3	19
Propane	0.048	0.016	16.5	0.38		2.2	
Propene	0.17	0.1	58.4	2.4		7.9	10
Acetylene	0.16	0.16	54.9	3.8		7.6	11.4
<i>i</i> -Butane	0.004	0.001	1.37	0.02		0.18	
<i>n</i> -Butane	0.013	0.0048	4.46	0.11		0.60	
t-2-Butene	0.01	0.0038	3.43	0.09		0.46	
1-Butene	0.034	0.019	11.0/	0.45		1.59	
<i>c</i> -2-Butene	0.0072	0.0028	2.47	0.07		0.33	
<i>i</i> -Pentane	0.0011	0.00081	0.38	0.02		0.05	
1.2 Dutediana	0.0041	0.0019	1.41	0.04		0.19	
2 Mothyl 1 bytono	0.039	0.027	13.4	0.05		1.65	
5-Wethyl-1-Dutelle	0.0038	0.0022	1.31	0.03		0.18	
2 Methyl 2 butene	0.0033	0.0013	1.15	0.04		0.15	
2-Methyl-1-butene	0.004	0.0015	1.57	0.03		0.18	
c-2-Pentene	0.0019	0.00094	0.65	0.07		0.09	
<i>n</i> -Hexane	0.0101	0.0071	3 47	0.17		0.07	
Isoprene	0.053	0.036	18.2	0.85		2.49	
2-Methyl-1-pentene	0.0028	0.0003	0.96	0.01		0.13	
Heptane	0.0056	0.004	1.92	0.09		0.26	
Benzene	0.13	0.089	44.6	2.1		6.1	
Toluene	0.15	0.11	51.5	2.6		7.0	
Formaldehyde	0.75	0.39	258	9.2	1,100	35	26
Methanol	1	0.32	343	7.5	3,800	46	57
Acetic acid	2.2	0.8	756	18.8	4,200	102	78
Formic acid	0.57	0.43	196	10.1	2,100	27	5
Total particulate matter	13	4.8	4,464	113	26,200	602	
Total particulate carbon	3.5	0.94	1,202	22	11,700	161	
Organic carbon	3	0.7	1,030	16.5	10,600	138	
Black carbon	0.47	0.24	161	5.6	1,500	22	
Chloride	0.26	0.0093	89.3	0.22		11.9	
Nitrate	0.19	0.05	65.3	1.2		8.8	
Sulfate	0.14	0.09	48.1	2.1		6.6	
Potassium	0.38	0.12	131	2.8	1,090	17.5	
CN ^e	5.9×10^{10}	1.7×10^{10}	2.0×10^{28}	4.0×10^{26}	1.1×10^{31}	2.7×10^{27}	

Table 4. Emission Factors for 49 Trace Gas and Particle Species for Miombo Woodland and Dambo Grassland Fires in Zambia and Emissions of These Species for Woodland and Grassland Savanna Fires in Zambia and Southern Africa^a

^aEmission factors are given in g kg^{-1} fuel burned; emissions are given in Gg. See section 5.5 for discussion.

^bSource is this paper.

^cSource is Andreae and Merlet [2001].

^dSource is *Bertschi et al.* [2003a].

^eUnits of CN emission factors and CN emissions are number of particles kg⁻¹ fuel burned and number of particles, respectively.

centages are $\sim 24.2\%$, 18.4%, 9.8%, and 9.2%, respectively. Percentages for the methyl halides CH₃Br, CH₃Cl, and CH₃I are $\sim 2.8\%$, 4.3%, and 0.5%, respectively.

[33] In Table 4 our estimate of HCN emissions from woodland and grassland savanna fires in southern Africa (~134 Gg) exceeds *Andreae and Merlet*'s [2001] estimate

of 90 Gg from global annual savanna burning. However, Andreae and Merlet calculated global annual emissions of HCN from savanna burning using an HCN emission factor of 0.025-0.031 g kg⁻¹ of fuel burned. If the global annual emissions of HCN from savanna burning are recalculated using the HCN emission factor derived from our measure-



Figure 6. Monthly biomass combusted by woodland and grassland savanna fires in southern Africa during the 2000 dry season.

ments $(0.31-0.37 \text{ g kg}^{-1} \text{ fuel burned})$, the total HCN emitted by global annual savanna fires is ~1.1 Tg. In a study of global sources and sinks of HCN, *Li et al.* [2000] estimated an annual source of 2.7–5.6 Tg from all types of biomass burning worldwide.

[34] In Table 4 we give estimates of emissions from woodland and grassland savanna fires in Zambia alone from May-October of 2000. Bertschi et al. [2003a] estimated annual savanna fire emissions from Zambia using emission factors from Yokelson et al. [2003] and model estimates of annual biomass burned from Hao and Liu [1994]. Bertschi et al. [2003a] also estimated annual emissions due to the use of biofuels in Zambia; they found that biofuel emissions are comparable to or exceed those from savanna burning for a number of species, including CH₄, CO, ethane, acetylene, acetic acid, formaldehyde, methanol, and NH₃. Annual biofuel emissions from Zambia reported by Bertschi et al. are shown in Table 4. When we compare them with our estimates of dry season emissions from woodland and grassland savanna fires in Zambia (Table 4), we find that the annual emissions of CH₄, ethane, ethene, acetylene, propene, formaldehyde, methanol, and acetic acid from biofuels are comparable to or exceed dry season emissions from woodland and grassland savanna fires in Zambia.

[35] A similar comparison can be made for the southern African region. Averaging two recent estimates of annual biofuel use in southern Africa [Yevich and Logan, 2003; Food and Agriculture Organization, 1999], we estimate that each year ~ 108 Tg of wood fuel is combusted along with \sim 4 Tg of charcoal. Combining these estimates with biofuel emission factors measured by Bertschi et al. [2003a], we estimate annual biofuel emissions (in Gg) in southern Africa of CO₂, CO, CH₄, NO_x, ethane, ethene, acetylene, acetic acid, formic acid, formaldehyde, and methanol to be $\sim 182,000, 12,300, 1360, 230, 214, 275,$ 182, 1020, 81, 398, and 569, respectively. The fourth and fifth columns of Table 4 show emissions of these species from woodland and grassland savanna fires, respectively, during the dry season of 2000 in southern Africa. Nearly three times as much CH4 is produced by annual biofuel use in southern Africa than by dry season woodland and grassland savanna fires in southern Africa. The relative importance of biofuel emissions of the other species is smaller although still substantial.

[36] Our emission estimates (Table 4) are subject to considerable uncertainty, most of which is due to uncertainty in the total area burned by woodland and grassland savanna fires in southern Africa during the 2000 dry season. In addition to the SPOT satellite burned-area retrievals used in this paper, preliminary burned area retrievals have been derived from measurements made by the Moderate-Resolution Imaging Spectroradiometer (MODIS) on the Terra satellite. In a comparison for September 2000 the burned area derived from MODIS measurements for southern Africa by S. Korontzi et al. (Modeling and sensitivity analysis of fire emissions in southern Africa during SAFARI 2000, submitted to Remote Sensing of Environment, 2003) is 55% larger than the SPOT burned area used in the present paper. The values used in the present paper for the available aboveground biomass fuel for miombo woodland and dambo grassland plots are uncertain to $\sim 30\%$. The combustion factors used in the present paper are uncertain to $\sim 30\%$, and the emission factors are uncertain to $\sim 10\%$. Therefore the emission estimates in Table 4 for woodland and grassland savanna fires in southern Africa during the 2000 dry season are uncertain to $\sim 70\%$.

[37] An additional source of uncertainty is emissions from residual smoldering combustion (RSC), which is biomass combustion that produces smoke that is not lofted by strong fire-induced convection. Emissions from RSC are generally not measured from aircraft, and they are not included in the emission factors reported in this paper. Residual smoldering combustion can occur in miombo woodland fires because of the presence of dead/downed logs in the fuel complex. Bertschi et al. [2003b] made the first measurements of emission factors from RSC in a miombo woodland. They found that the inclusion of RSC emissions in estimating total emissions from miombo fires had little effect on many compounds but had a significant effect on some. Little data exist on the amount and composition of RSC smoke. Therefore we will not explore the topic further here except to note that some additional uncertainty due to RSC emissions may occur for some compounds, particularly methane and methanol.

6. Conclusions

[38] In this paper we have described airborne measurements of the emissions and initial evolution of trace gases and particles from a miombo woodland and a dambo grassland fire in Zambia. The principal conclusions from analyses of these measurements are as follows.

[39] 1. Woodland and grassland savanna fires in southern Africa during the 2000 dry season were estimated to account for ~12.3%, 12.6%, 5.9%, 10.3%, 4.0%, and 7.5% of annual emissions of CO₂, CO, total hydrocarbons, NO_x (as NO), NH₃, and SO₂, respectively, from all types of savanna fires worldwide. For total particulate matter, black carbon, organic carbon, and potassium the corresponding percentages are ~17.5%, 11.1%, 9.9%, and 12.2%, respectively. For oxygenated species such as formaldehyde, acetic acid, formic acid, and methanol the percentages are ~24.2%, 18.4%, 9.8%, and 9.2%, respectively. For CH₃Br, CH₃Cl, and CH₃I they are ~2.8%, 4.3%, and 0.5%, respectively. Our emissions estimates for woodland and

grassland savanna fires during the 2000 dry season in southern Africa are uncertain to \sim 70%.

[40] 2. Average annual emissions of CH_4 , ethane, ethene, acetylene, propene, formaldehyde, methanol, and acetic acid from the use of biofuels in Zambia are comparable to or exceed dry season emissions from miombo woodland and dambo grassland fires in Zambia. Average annual emissions of methane from the use of biofuels in southern Africa are nearly 3 times higher than those from woodland and grassland savanna fires in southern Africa during the dry season of 2000.

[41] Acknowledgments. We thank all members of the UW-CARG and the pilots of the UW Convair-580 for their valuable help in obtaining measurements. We thank Dan Jaffe for help in calibrating the gas instruments. This study was carried out as part of the SAFARI 2000 Southern African Regional Science Initiative. Research support from grants NAG5-9022 and NAG5-7675 from NASA's Radiation Science Program and grants ATM-9901624 and ATM-9900494 from NSF's Division of Atmospheric Sciences is gratefully acknowledged.

References

- Andreae, M. O., and P. Merlet (2001), Emission of trace gases and aerosols from biomass burning, *Global Biogeochem. Cycles*, 15, 955–966.
- Annegarn, H. J., L. Otter, R. J. Swap, and R. J. Scholes (2002), Southern Africa's ecosystem in a test-tube: A perspective on the Southern African Regional Science Initiative (SAFARI 2000), S. Afr. J. Sci., 98, 111–113.
- Bertschi, I. T., R. J. Yokelson, D. E. Ward, T. J. Christian, and W. M. Hao (2003a), Trace gas emissions from the production and use of domestic biofuels in Zambia measured by open-path Fourier transform infrared spectroscopy, J. Geophys. Res., 108(D13), 8469, doi:10.1029/ 2002JD002158.
- Bertschi, I., R. J. Yokelson, D. E. Ward, R. E. Babbitt, R. A. Susott, J. G. Goode, and W. M. Hao (2003b), Trace gas and particle emissions from fires in large diameter and belowground biomass fuels, *J. Geophys. Res.*, 108(D13), 8472, doi:10.1029/2002JD002100.
- Bodhaine, B. A. (1995), Aerosol adsorption measurements at Barrow, Mauna Loa, and the south pole, *J. Geophys. Res.*, 100, 8967–8975.
- Colman, J. J., A. L. Swanson, S. Meinardi, B. C. Sive, D. R. Blake, and F. S. Rowland (2001), Description of the analysis of a wide range of volatile organic compounds in whole air samples collected during PEM-Tropics A and B, *Anal. Chem.*, 73, 3723–3731.
- Crutzen, P. J., and M. O. Andreae (1990), Biomass burning in the tropics: Impact on atmospheric chemistry and biogeochemical cycles, *Science*, 250, 1669–1678.
- International Geosphere-Biosphere Programme (1997), The miombo network: Framework for a terrestrial transect study of land-use and landcover change in the miombo ecosystems of central Africa, edited by P. V. Desanker et al., *IGBP Rep. 41*, Stockholm.
- Duncan, B. N., R. V. Martin, A. C. Staudt, R. Yevich, and J. A. Logan (2003), Interannual and seasonal variability of biomass burning emissions constrained by satellite observations, *J. Geophys. Res.*, 108(D2), 4100, doi:10.1029/2002JD002378.
- Eck, T. F., et al. (2003), Variability of biomass burning aerosol optical characteristics in southern Africa during the SAFARI 2000 dry season campaign and a comparison of single scattering albedo estimates from radiometric measurements, *J. Geophys. Res.*, 108(D13), 8477, doi:10.1029/2002JD002321.
- Fishman, J., K. Fakhruzzaman, B. Cros, and D. Nganga (1991), Identification of widespread pollution in the southern hemisphere deduced from satellite analyses, *Science*, 252, 1693–1696.
- Food and Agriculture Organization (1999), The role of wood energy in Africa, FAO Working Pap. FOPW/99/3, Rome.
- Gao, S., D. A. Hegg, P. V. Hobbs, T. W. Kirchstetter, B. I. Magi, and M. Sadilek (2003), Water-soluble organic components in aerosols associated with savanna fires in southern Africa: Identification, evolution, and distribution, *J. Geophys. Res.*, 108(D13), 8491, doi:10.1029/2002JD002324.
- Gundel, L. A., R. L. Dod, H. Rosen, and T. Novakov (1984), The relationship between optical attenuation and black carbon concentrations for ambient and source particles, *Sci. Total Environ.*, 36, 197–202.
- Hao, W. M., and M.-H. Liu (1994), Spatial and temporal distribution of tropical biomass burning, *Global Biogeochem. Cycles*, 8, 495–503.
- Haywood, J. M., S. R. Osborne, P. N. Francis, A. Keil, P. Formenti, M. O. Andreae, and P. H. Kaye (2003), The mean physical and optical properties of regional haze dominated by biomass burning aerosol measured

from the C-130 aircraft during SAFARI 2000, J. Geophys. Res., 108(D13), 8473, doi:10.1029/2002JD002226.

- Hobbs, P. V., P. Sinha, R. J. Yokelson, T. J. Christian, D. R. Blake, S. Gao, T. W. Kirchstetter, T. Novakov, and P. Pilewskie (2003), Evolution of gases and particles from a savanna fire in South Africa, *J. Geophys. Res.*, 108(D13), 8485, doi:10.1029/2002JD002352.
- Hoffa, E. A., D. E. Ward, W. M. Hao, R. A. Susott, and R. H. Wakimoto (1999), Seasonality of carbon emissions from biomass burning in a Zambian savanna, *J. Geophys. Res.*, 104, 13,841–13,853.
- Kirchstetter, T. W., T. Novakov, P. V. Hobbs, and B. Magi (2003), Airborne measurements of carbonaceous aerosols in southern Africa during the dry biomass burning season, *J. Geophys. Res.*, 108(D13), 8476, doi:10.1029/ 2002JD002171.
- Li, Q., D. J. Jacob, I. Bey, R. M. Yantosca, Y. Zhao, Y. Kondo, and J. Notholt (2000), Atmospheric hydrogen cyanide (HCN): Biomass burning source, ocean sink?, *Geophys. Res. Lett.*, 27, 357–360.
- Magi, B. I., and P. V. Hobbs (2003), Effects of humidity on aerosols in southern Africa during the biomass burning season, *J. Geophys. Res.*, 108(D13), 8495, doi:10.1029/2002JD002144.
- Novakov, T. (1981), Microchemical characterization of aerosols, in *Nature, Aim, and Methods of Microchemistry*, edited by H. Malissa, M. Grasserbauer, and R. Belcher, pp. 141–165, Springer-Verlag, New York.
- Novakov, T. (1982), Soot in the atmosphere, in *Particulate Carbon: Atmospheric Life Cycle*, edited by G. T. Wolff and R. L. Klimisch, pp. 19–41, Plenum, New York.
- Radke, L. F., J. L. Stith, D. A. Hegg, and P. V. Hobbs (1978), Airborne studies of particles and gases from forest fires, *J. Air Pollut. Control Assoc.*, 28, 30–34.
- Radke, L. F., D. A. Hegg, J. H. Lyons, C. A. Brock, and P. V. Hobbs (1988), Airborne measurements on smokes from biomass burning, in *Aerosols and Climate*, edited by P. V. Hobbs and M. P. McCormick, pp. 411–422, A. Deepak, Hampton, Va.
- Reid, J. S., P. V. Hobbs, R. J. Ferek, D. R. Blake, J. V. Martins, M. R. Dunlap, and C. Liousse (1998), Physical, chemical, and optical properties of regional hazes dominated by smoke in Brazil, *J. Geophys. Res.*, 103, 32,059–32,080.
- Rosen, H., and T. Novakov (1983), Optical-transmission through aerosol deposits on diffusely reflective filters—A method for measuring the absorbing component of aerosol-particles, *Appl. Opt.*, 22, 1265–1267.
- Scholes, R. J., J. Kendall, and C. O. Justice (1996), The quantity of biomass burned in southern Africa, J. Geophys. Res., 101, 23,667–23,676.
- Shea, R. W., B. W. Shea, J. B. Kauffman, D. E. Ward, C. I. Haskins, and M. C. Scholes (1996), Fuel biomass and combustion factors associated with fires in savanna ecosystems of South Africa and Zambia, J. Geophys. Res., 101, 23,551–23,568.
- Silva, J. M. N., J. M. C. Pereira, A. I. Cabral, A. C. L. Sá, M. J. P. Vasconcelos, B. Mota, and J.-M. Grégoire (2003), An estimate of the area burned in southern Africa during the 2000 dry season using SPOT-VEGETATION satellite data, J. Geophys. Res., 108(D13), 8498, doi:10.1029/2002JD002320.
- Sinha, P., P. V. Hobbs, R. J. Yokelson, I. T. Bertschi, D. R. Blake, I. J. Simpson, S. Gao, T. W. Kirchstetter, and T. Novakov (2003a), Emissions of trace gases and particles from savanna fires in southern Africa, *J. Geophys. Res.*, 108(D13), 8487, doi:10.1029/2002JD002325.
- Sinha, P., P. V. Hobbs, R. J. Yokelson, D. R. Blake, S. Gao, and T. W. Kirchstetter (2003b), Distributions of trace gases and aerosols during the dry biomass burning season in southern Africa, *J. Geophys. Res.*, 108(D17), 4536, doi:10.1029/2003JD003691.
- Susott, R. A., G. J. Olbu, S. P. Baker, D. E. Ward, J. B. Kauffmann, and R. W. Shea (1996), Carbon, hydrogen, nitrogen, and thermogravimetric analysis of tropical ecosystem biomass, in *Biomass Burning and Global Change*, vol. 1, edited by J. S. Levine, pp. 249–259, MIT Press, Cambridge, Mass. Swap, R. J., H. J. Annegarn, J. T. Suttles, M. D. King, S. Platnick, J. L.
- Swap, R. J., H. J. Annegarn, J. T. Suttles, M. D. King, S. Platnick, J. L. Privette, and R. J. Scholes (2003), Africa burning: A thematic analysis of the Southern African Regional Science Initiative (SAFARI 2000), *J. Geophys. Res.*, 108(D13), 8465, doi:10.1029/2003JD003747.
- Ward, D. E., and W. M. Hao (1992), Air toxic emissions from burning of biomass globally—Preliminary estimates, paper presented at 85th Annual Meeting and Exhibition, Air and Waste Manage. Assoc., Vancouver, B. C., Canada.
- Ward, D. E., and C. C. Hardy (1991), Smoke emissions from wildland fires, *Environ. Int.*, 17, 117–134.
- White, F. (1981), UNESCO/AETFAT/UNSO vegetation map of Africa, scale 1:5,000,000, UNESCO, Paris.
- White, F. (1983), Vegetation of Africa: A Descriptive Memoir to Accompany the UNESCO/AETFAT/UNSO Vegetation Map of Africa, UNESCO, Paris.
- Yevich, R., and J. A. Logan (2003), An assessment of biofuel use and burning of agricultural waste in the developing world, *Global Biogeochem. Cycles*, 17(4), 1095, doi:10.1029/2002GB001952.

Yokelson, R. J., I. T. Bertschi, T. J. Christian, P. V. Hobbs, D. E. Ward, and W. M. Hao (2003), Trace gas measurements in nascent, aged, and cloudprocessed smoke from African savanna fires by airborne Fourier transform infrared spectroscopy (AFTIR), J. Geophys. Res., 108(D13), 8478, doi:10.1029/2002JD002322. S. Gao, Department of Environmental Science and Engineering, California Institute of Technology, Pasadena, CA 91125, USA. (sgao@its. caltech.edu)

P. V. Hobbs and P. Sinha, Department of Atmospheric Sciences, University of Washington, Box 351640, Seattle, WA 98195-1640, USA. (phobbs@atmos.washington.edu; psinha@atmos.washington.edu)

T. W. Kirchstetter, Environmental Energy Technologies Division, Lawrence Berkeley National Laboratory, Berkeley, CA 94720, USA. (twkirchstetter@lbl.gov)

R. J. Yokelson, Department of Chemistry, University of Montana, Missoula, MT 59812, USA. (byok@selway.umt.edu)

D. R. Blake, Department of Chemistry, 516 Rowland Hall, University of California, Irvine, Irvine, CA 92697-2025, USA. (drblake@uci.edu)